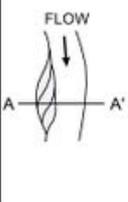
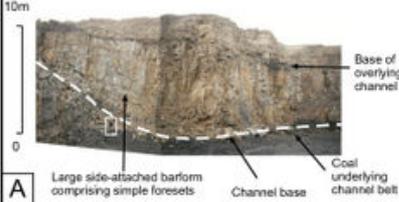
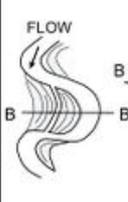
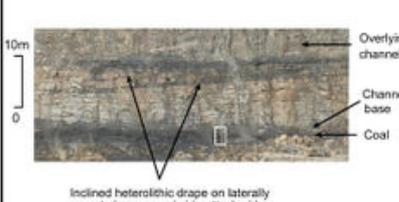
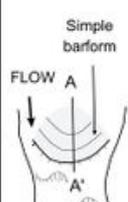
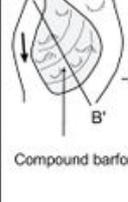


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| Type | Description | Interpretation | Sketch | Examples |
|---------------|--|---|--|---|
| SIDE-ATTACHED | SIMPLE: Thick, broadly lenticular units comprising sandy, high-angle (10-25 deg) foresets attached to the margins of a channel. Up to 11m in height, total width up to 100 m. Rare internal cross-beds. Foresets up to 0.5m thick. Angular to asymptotic foreset bases. Rare chute channels occur, incised into the top. Foreset dip in same direction as foresets from adjacent sedimentary structures | Downstream accretion of sand onto the lee side of a large, side-attached, slip-face fronted bar. Avalanching and grainfall processes important over the crest of the bar. Modified by erosion during falling stage. Analogous to the alternate bars of Miall (1966) and have been recognised from the modern day Platte River of Colorado (Crowley 1983) and the Jurassic Scalby Formation (Eschard et al. 1991). |  <p>Simple barform Channel margin (N.B. Barform accretes in a downstream direction)</p> |  <p>10m A Large side-attached barform comprising simple foresets Channel base Coal underlying channel belt Base of overlying channel</p> |
| | COMPOUND: Heterolithic barform comprising low-angle (<10 deg.) 3rd order bounding surfaces that dip perpendicular to the channel axis. Internally cross-bedded & cross-laminated. Bedding is asymptotic or sigmoidal and typically dips normal to palaeocurrent trend. Individual beds show vertical & lateral upwards fining and decrease in scale of sedimentary structures through the bed. Upper parts are typically mudier and scoured into by small channels. From 4 to 15m in thickness. Individual beds up to 100m in length. Barform deposit up to 400m in width | Lateral accretion of dunes and ripples onto a gently dipping, side-attached (point) bar that dips perpendicular to main channel trend. Fining produced by upwards diminution in flow velocities up the bar. Scours across the bar top represent chute channels cut during high stage. Well documented from modern and ancient rivers. Represents deposition on the point bar of a meandering channel (see Leopold & Wolman 1960, Jackson 1976, Nijman & Puigdefabrigas 1978, Smith 1987). |  <p>Chute channel Channel margin Compound barform (N.B. Barform accretes laterally with respect to channel trend)</p> |  <p>10m B Overlying channel Channel base Coal Inclined heterolithic drape on laterally accreted compound side-attached bar</p> |
| MID-CHANNEL | SIMPLE: Thick, broadly lenticular units comprising sandy, high-angle (10-25 deg.) foresets present within the central parts of a channel. Rare internal cross-beds. Foresets up to 0.5m thick. Angular to asymptotic foreset bases. Tend to flatten out downstream and may fill deeper scours along channel bases. Up to 7m thick. Barform produces a deposit up to 50m wide and > 50m long | Downstream-accretion of sand onto the lee side of a large, mid-channel, slip-face fronted bar. Avalanching and grainfall processes important over the bar crest. Modified by falling stage processes. Documented from the Triassic Hawkesbury Sandstone of Australia (Conaghan & Jones 1975, Rust & Jones 1987). They are up to 15m in height in the Brahmaputra (Coleman 1969, Bristow 1987, 1993). |  <p>Simple barform 'Simple' foresets (N.B. Barform accretes in a downstream direction)</p> |  <p>10m C Erosive channel base Foresets of simple downstream accreting mid-channel bar</p> |
| | COMPOUND: Thick, sandy units comprising laterally persistent bounding surfaces that dip at low angles (< 10 deg.) down-palaeocurrent. Form within the central parts of a channel. Internally cross-bedded, with cross-bed orientations parallel to the dip of the barform bounding surfaces. Rare reactivation surfaces. Downcurrent-descending cross-beds. Up to 12m thick, laterally persistent for 70m in width | Large low-relief barform; lacks a slipface & has a low-angle lee that merges downstream & laterally with thalweg deposits. The barform is built up of downstream accreting superimposed dunes and ripples. Equivalent to the downstream-accretion macroforms of Miall (1988a, b). Reported in modern sandy braided rivers, e.g. the Tana (Collinson 1970), South Saskatchewan (Cant & Walker 1978) & Brahmaputra (Bristow 1987, 1993). |  <p>Dunes migrate down low angle barform front Bounding surface Internal sets of cross-bedding Compound barform (N.B. Barform accretes in a downstream direction)</p> |  <p>8m D Inclined barform front accretion surfaces</p> |

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Summary

| | |
|--------------------|--|
| description | <p>English: Table 2 Main types of barform channel elements recognized from the Upper Carboniferous succession in northern Germany.</p> <p>From: Carboniferous hydrocarbon resources: the southern North Sea and surrounding onshore areas, edited by J. D. Collinson, D. J. Evans, D. W. Holliday, N. S. Jones. Published as volume 7 in the Occasional Publications series of the Yorkshire Geological Society, Copyright Yorkshire Geological Society 2005.</p> |
| source | Yorkshire Geological Society |
| author | Neil S. Jones and Brian W. Glover |

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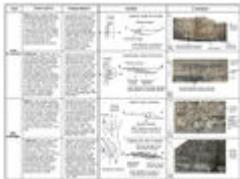
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